Vt 2011

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Introduction to hydrological modelling and the HBV model









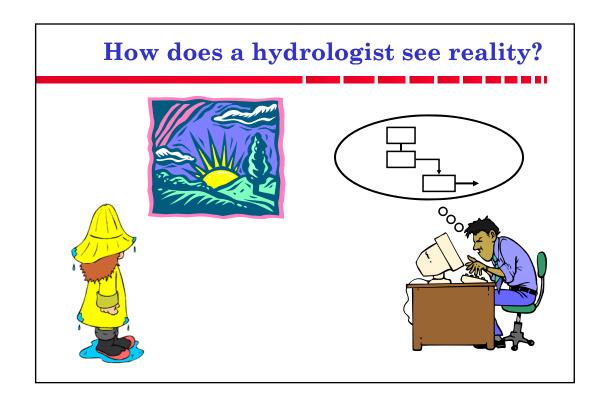


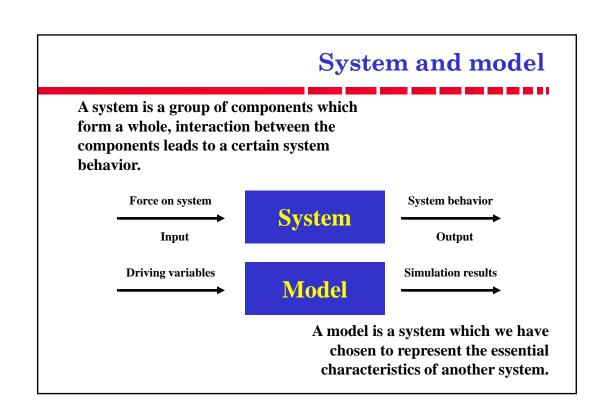


How do we see reality?



Figur 3.2. Vi väljer och omtolkar den information som vi sedan baserar våra subjektiva tankemodeller på. (Av Nils Peterson, ur Forskning och Framsteg.)





Why use a model?

- ◆Forecast / Prediction
- ♦ Instead of expensive and time consuming measurements
- 'What happens if' questions
- Summarize and test knowledge
- Education



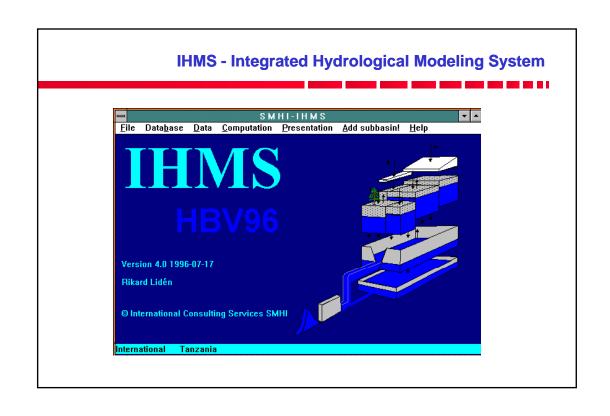
Classification of models

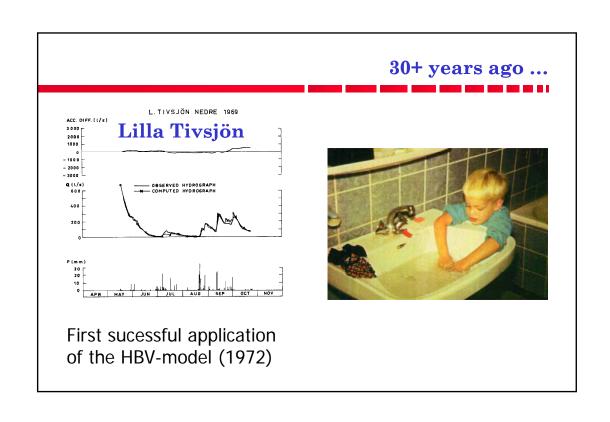
- ◆Symbolic (mathematic) scale-/ analog model
- ◆Empirical conceptual physical
- ◆Linear non-linear
- **(Lumped)** distributed
- Deterministic stochastic
- ♦ Static dynamic
- ◆Time-variant time-invariant
- **♦**Event based -continuous
- **◆**Complete partial
- ◆General special purpose

The HBV model ...

- ... has been developed by Sten Bergström
- ... is a conceptual model for runoff simulation
- ... has a simple structure
- ... is semi-distributed, i.e., allows to divide the catchment into subbasins, elevation and vegetation zones
- ... is easy to understand, learn and apply
- ... has been applied to many catchments in Sweden and abroad
- ... provided good results in most applications
- ... has become a standard tool for runoff studies in the Nordic countries
- ... needs a moderate amount of input data
- ... can be run on a PC (286 or better)
- ... exists in different versions (Swedish, Norwegian, Finnish, Swiss, ...)
- ... (partly) used in other models







30+ years with the HBV-model - Lilla Tivsjön 2002



... whereas the HBV-model is still in pretty good shape

Input data

- Areal precipitation
 - Weighted mean
 - Elevation zones (lapse rate ~10% per 100 m)
- Temperature
 - Weighted mean
 - Elevation zones (lapse rate 0.6 °C per 100 m)
- Potential evaporation
 - Penman formula or measurements
 - Usually long-term monthly mean values

Input data (cont.)

Correction of long-term potential evaporation

$$\begin{aligned} \mathsf{E}_{\mathsf{pot}}(t) &= (1 + C_{\mathsf{ET}}(\mathsf{T}(t) - \mathsf{T}_{\mathsf{M}})) \; \mathsf{E}_{\mathsf{pot},\mathsf{M}} \\ & (\mathsf{BUT:} \; \; 2 \; \mathsf{E}_{\mathsf{pot},\mathsf{M}} \geq \mathsf{E}_{\mathsf{pot}}(t) \geq 0) \end{aligned}$$

E_{pot}(t) potential evaporation at day t [mm d⁻¹]

C_{ET} correction factor [°C⁻¹]

T(t) temperature at day t [°C]

T_M long-term mean temperature for this day of the year [°C]

E_{pot,M} long-term mean evaporation for this day of the year [mm d⁻¹]

Model structure Input data Routine Output data Precipitation, Snow Snow pack, routine temperature snow-melt Soil Act. routine evaporation, evaporation, precipitation, 'soil moisture', snowmelt groundwater recharge Response Runoff to Groundwater PRECIPITATION function recharge, stream, (pot. 'Groundwater evaporation) level' Snow routine Routing Runoff to Runoff at outlet routine stream Soil moisture routine Response function Routing routine RUNOFF

Snow routine

- Accumulation of precipitation as snow if temperature
- Degree-day method for snowmelt (degree-day factor C_{FMAX} [mm d⁻¹ °C ⁻¹])

$$M=C_{FMAX}(T-T_T)$$
 [mm d⁻¹]



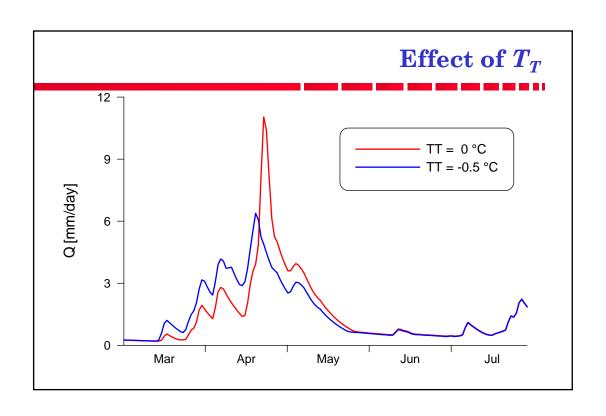
C_{FMAX} typically around 4 mm d⁻¹ °C ⁻¹, lower values for forested areas compared to open areas

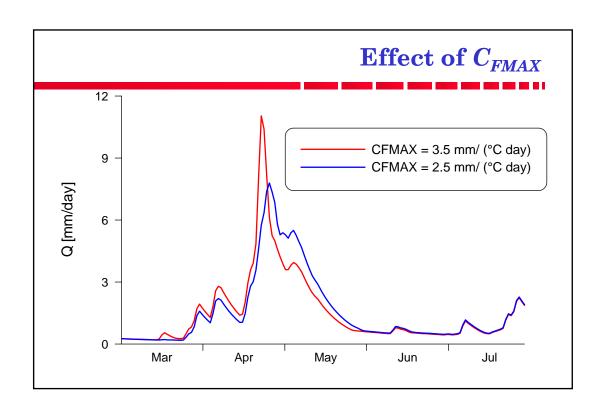
Snow routine (cont.)

- Snow pack retains melt water until amount exceeds a certain portion (C_{WH}, usually 0.1) of the water equivalent of the snow pack.
- When temperatures decreases below T_T water refreezes again

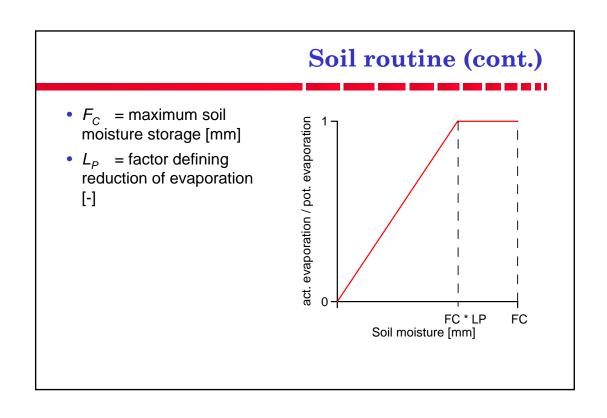
$$M = C_{FR} C_{FMAX} (T_T - T)$$
 [mm h⁻¹]
 $C_{FR} = \sim 0.05$ [-]

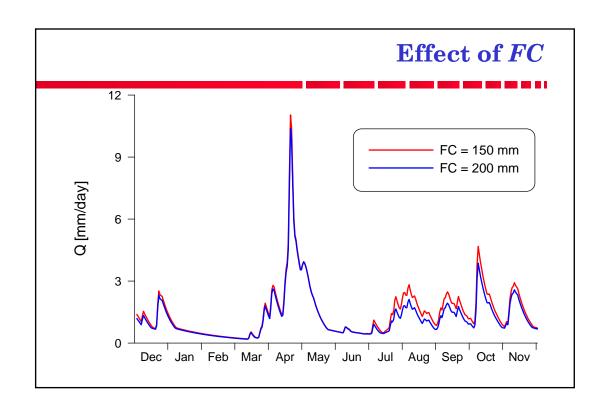
- All precipitation which is simulated to be snow is multiplied by a correction factor, S_{CF}[-]
- Calculations are carried out separately for each elevation and vegetation zone

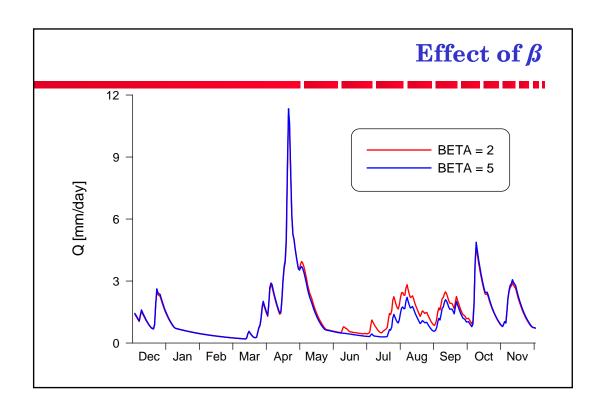




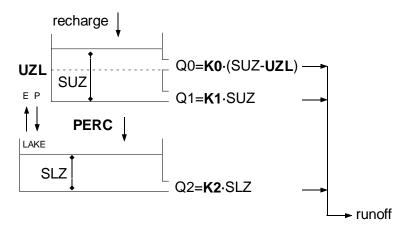
• F_C maximum soil moisture storage [mm] • ß shape parameter [-] $\frac{\text{recharge}}{P} = \left(\frac{S_{sm}}{F_{C}}\right)^{\beta}$ $\frac{\text{groundwater recharge}}{Soil moisture [mm]}$ • Soil moisture [mm]





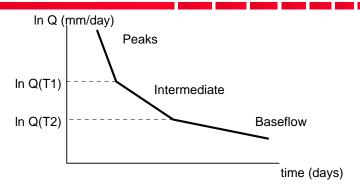


Response function



- 5 parameter, two boxes
- · Upper box: shallow ground water, lower box: deeper groundwater

Response function (cont.)



• Slope of the recession:

Peaks:

 $K_0 + K_1 + K_2$

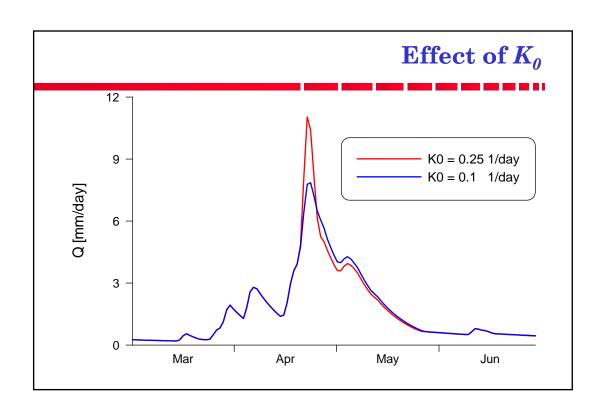
Intermediate: K₁+ K₂

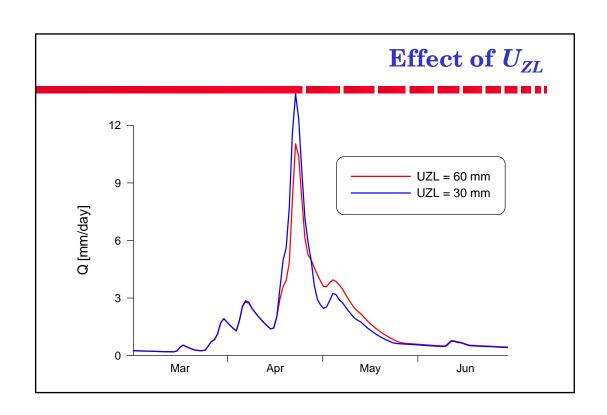
Baseflow:

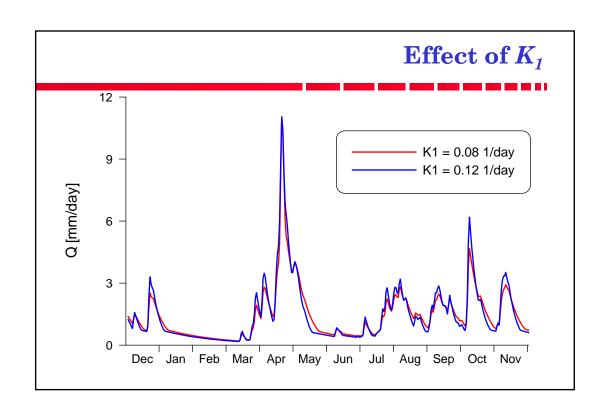
Thresholds:

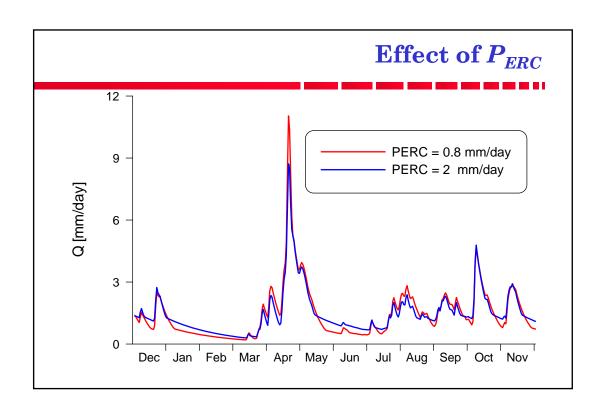
$$\mathsf{Q}(\mathsf{T}_1) = \sim \mathsf{P}_{\mathsf{ERC}} {+} \mathsf{K}_1 \mathsf{U}_{\mathsf{ZL}}$$

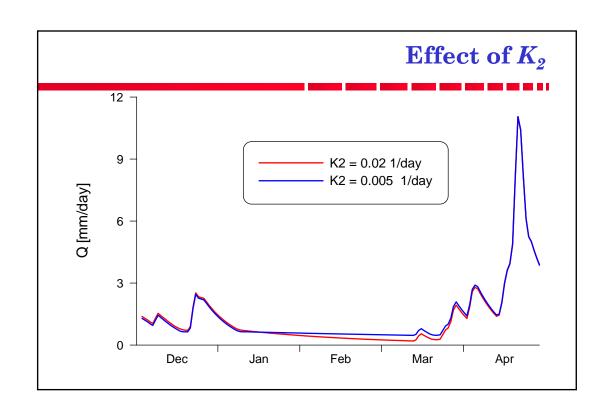
 $Q(T_2) = \sim P_{ERC}$

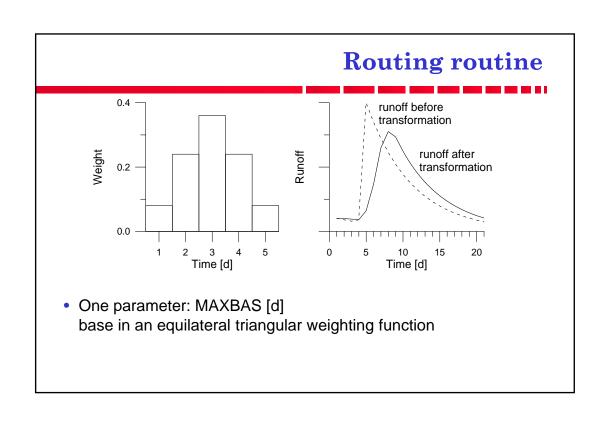


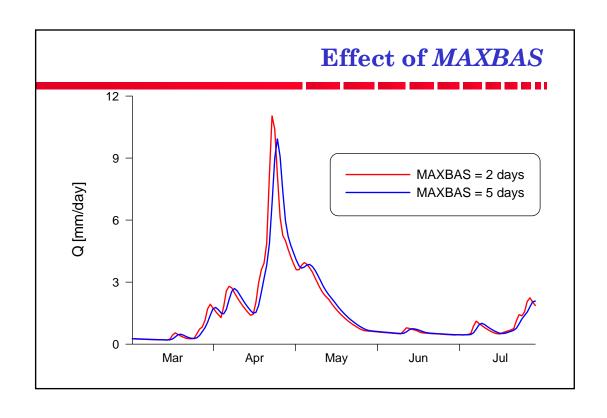


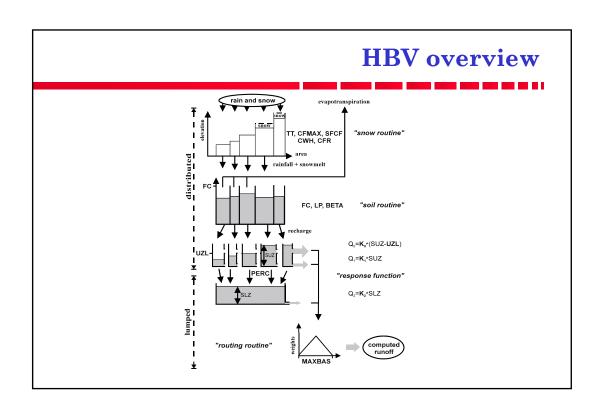












Calibration of the HBV model

- · Trial-and-error or automatic calibration
- Different criteria can be used to assess the fit of simulated runoff to observed runoff:

visual inspection of plots with \mathbf{Q}_{Sim} and \mathbf{Q}_{Obs} statistical criteria



 The coefficient of efficiency, R_{eff}, is normally used for assessment of simulations by the HBV model:

$$R_{eff} = 1 - \frac{\sum (Q_{Sim}(t) - Q_{Obs}(t))^{2}}{\sum (Q_{Obs}(t) - \overline{Q}_{Obs})^{2}}$$

 R_{eff} = 1 -> Perfect fit; R_{eff} = 0 -> Simulation as good (or poor) as the constantvalue prediction, R_{eff} < 0 -> Very poor fit

R_{eff} not equal r²!!!

Calibration of the HBV model (cont.)

- Calibration period should include a variety of hydrological events
- Normally 5 to 10 years sufficient to calibrate the model
- Split-sample test: test of model performance with calibrated parameters for an independent period
- Different objective functions (e.g., efficiency based on log Q, or 'punishment' for volume error)
- Problems

Parameter uncertainty
Internal model consistency

Parameter values for ungauged catchments

Applications of the HBV model

The HBV model is/can be used ...

- ... to extend runoff data series (or filling gaps)
- ... for data quality control
- ... for water balance studies
- ... for runoff forecasting (flood warning and reservoir operation)
- ... to compute design floods for dam safety
- ... to investigate the effects of changes within the catchment
- ... to simulate discharge from ungauged catchments
- ... to simulate climate change effects

